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ABSTRACT

Peatland development and carbon accumulation on the Pacific coast of Canada have received little attention in paleoecological studies despite wetlands being common landscape features. Here, we present a multi-proxy paleoenvironmental study of an ombrotrophic bog in coastal British Columbia. Following decreases in relative sea level, the wetland was isolated from marine waters by 13,300 cal yr BP. Peat composition, non-pollen palynomorph, and C and N analyses demonstrate terrestrialization from an oligotrophic lake to a marsh by 11,600 cal yr BP, followed by development of a poor fen, and then a drier ombrotrophic bog by 8700 cal yr BP. Maximum carbon accumulation occurred during the early Holocene fen stage, when seasonal differences in insolation were amplified. This highlights the importance of seasonality in constraining peatland carbon sequestration by enhancing productivity during summer and reducing decomposition during winter. Pollen analysis shows that *Pinus contorta* dominated regional forests by 14,000 cal yr BP. Warm and relatively dry summers in the early Holocene allowed *Pseudotsuga menziesii* to dominate lowland forests 11,200–7000 cal yr BP. *Tsuga heterophylla* and *P. menziesii* formed coniferous forest in the mid- and late Holocene. Tephra matching the mid-Holocene Glacier Peak-Dusty Creek assemblage provides evidence of its most northwesterly occurrence to date.

Keywords: peatlands; terrestrialization; carbon accumulation; nitrogen accumulation; pollen; non-pollen palynomorphs; plant macrofossils; coniferous forest; Glacier Peak tephra; coastal British Columbia

INTRODUCTION

Paleoecological studies along the north Pacific coast of North America have largely focussed on inferring vegetation change since the last glacial maximum through pollen analysis of lake sediments. This research has revealed a rich paleoecological history, marked by the early to mid-Holocene establishment of temperate rainforests with some of the planet's largest stores of above-ground biomass per unit area (Pan et al., 2013). Few studies have focussed on understanding peatland dynamics in this maritime setting, despite wetlands being common landscape features and important carbon (C) stores, and even fewer have inferred long-term rates of C accumulation. At a slope bog on the north coast of British Columbia (BC), Turunen and Turunen (2003) determined that peat accumulation began 12,000 cal yr BP via forest paludification and that mean C accumulation rates over the last 8500 cal yr were only 8.6 g C/m^2/cal yr (Loisel et al., 2014). Lacourse and Davies (2015) documented a higher mean rate (16.1 g C/m^2/cal yr) for the last 10,000 cal yr at a flat *Sphagnum* bog on northern Vancouver Island that formed through terrestrialization. These C accumulation rates are similar to rates at peatlands to the north on the south coast of Alaska (9–19 g C/m^2/cal yr: Jones and Yu, 2010; Loisel et al., 2014; Nichols et al., 2014), but lower than Loisel et al.'s (2014) estimate for northern peatlands generally (23 g C/m^2/cal yr) and significantly lower than those in some continental peatlands (~30 g C/m^2/cal yr; Yu et al., 2014; Zhao et al., 2014). Documenting long-term C accumulation rates, particularly in coastal BC where there are few studies, is important for understanding peatland C sequestration, improving estimates of Holocene C stocks, and clarifying the effects of climate change on peatlands and their role in global change as sinks and sources of carbon dioxide and methane (Loisel et al., 2017).

This paper focusses on the paleoecology and C accumulation of a wetland on the east coast of Vancouver Island, the largest island on the Pacific coast of North America. Vancouver Island is separated from the BC mainland by channels that range in width from as little as a few kilometers to as much as 55 km. The island is characterized by steep climatic and ecological gradients due primarily to the Vancouver Island Ranges that run the length of the island (Fig. 1), creating a rainshadow that is magnified further by the Olympic Mountains to the south in Washington. Mean annual precipitation exceeds 3000 mm on the north and west coasts of the island but is only 600 mm on the dry southeastern coast. Much of the BC coast supports closed-canopy coniferous rainforest with bog-forest complexes that are particularly abundant along the north coast. The narrow strip of lowlands on the southeast coast of Vancouver Is. (Fig. 1) are characterized by long, dry summers and relatively open *Pseudotsuga menziesii*-dominated forest.

Here, we use multiple paleoenvironmental proxies to infer the developmental history of an ombrotrophic bog on Vancouver Is. and changes in regional forest composition over the last 14,000 cal yr. We combine pollen, non-pollen palynomorphs, plant macrofossils and bulk chemical analyses, including carbon (C) and nitrogen (N) isotopes, to document changes in regional and local plant communities as well as hydrological and edaphic conditions. We also compare long-term rates of C accumulation to other peatland records and Holocene climate change. This study advances our understanding of wetland succession, long-term C accumulation and peatland dynamics in a temperate maritime setting. We also further refine the paleoecological history of coastal BC by providing a new pollen record of postglacial forest dynamics from an area of the coast that has received little attention in previous research.

STUDY SITE

Grant's Bog (49°47.3'N, 125°07.6'W, 80 m asl) is located 7 km from the coast in the Black Creek watershed of the eastern coastal lowlands of Vancouver Is., British Columbia (Fig. 1). The area supports coniferous forest dominated by *Pseudotsuga menziesii* and *Tsuga heterophylla*. Mean July temperature near the study site is 17.1°C, mean January temperature is 2.8°C, and the number of frost-free days is at least 280 (Black Creek weather station; Environment Canada, 2018). Mean annual precipitation is 1645 mm/yr; summers are generally dry with most precipitation falling as rain between October and March (Fig. 1).

Grant's Bog (informal name) is part of a 70-ha wetland complex that includes 7.5-ha of marsh along the southwestern margin that is covered by emergent *Nuphar polysepala* with a peripheral fringe of *Dulichium arundinaceum*, and a shallow, open-water pond (1.8 ha) in the southeastern corner. These shallow-water ecosystems occupy slightly deeper topographic depressions than the *Sphagnum*-ericad bog that characterizes most of the wetland complex. Plant cover in the bog is dominated by *Sphagnum* mosses (*S. fuscum*, *S. angustifolium*, *S. capillifolium, S. palustre*) and ericaceous shrubs (*Rhododendron groenlandicum*, *Kalmia microphylla* var. *occidentalis*, *Vaccinium uliginosum*). Other common species include *V. oxycoccus*, *Rubus chamaemorus*, *Eriophorum chamissonis*, *Rhynchospora alba* and *Drosera rotundifolia*. *Empetrum nigrum*, *Myrica gale* and stunted *Pinus contorta* var. *contorta* are present in low abundance. The water table in the bog was 16 cm below the surface at the coring location in July 2013 and mean water pH was 3.6. Golinski (2004) documented mean annual water table fluctuations of 35 cm.

METHODS

A 810-cm peat and sediment core was collected from Grant’s Bog in July 2013 using a 'Russian' D-corer with a 50 cm-long and 5 cm-diameter semi-cylindrical chamber. We alternated between two boreholes located 25 cm apart and collected sections with 10 cm of overlap. Nine AMS radiocarbon ages were obtained on plant macrofossils or organic lake sediment (Table 1). Six of these ages are at depths where a stratigraphic change occurs, which allows accumulation rates to be more reliably estimated than dating at systematic intervals. The IntCal13 dataset (Reimer et al., 2013) was used to calibrate ^{14}C ages to calendar years (cal yr BP). An age-depth model was built on calendar age probability distributions and an age of −63 cal yr BP for the top of the core, using 10,000 iterations of a smooth spline in the ‘clam’ package (Blaauw, 2010) in R (R Core Team, 2017). The age at 727 cm on wood was excluded from the model because it is out of stratigraphic order and considerably younger than the older ages immediately above and below. The 'Bacon' package (Blaauw and Christen, 2011) was not used to build a chronology because that approach produces a model that is more or less equivalent to linear interpolation but discounts the ^{14}C age at 626 cm, which provides important chronological control on the transition to a terrestrial environment.

Loss-on-ignition was conducted on 1–2 cm^3 samples taken at 2–4 cm intervals along the length of the core. Samples were dried at 105°C for 20 hr and then ignited at 550°C for 4 hr. C and N analyses were conducted at a resolution of <150 cal yr between samples in the peat portion of the core. Samples of 2–3 cm^3 were dried for 48 hr at 55°C and ground to a fine powder (<125 μm) with a Retsch MM 200 ball mill. Tin capsules (5×8 mm) were then packed with 3–5 mg of homogenized peat and analyzed on a Costech ECS 4010 thermal combustion EA coupled to a Thermo Finnigan DELTAplus Advantage IRMS. Replicate analyses were conducted on 15% of samples. Standards including acetanilide (71.09% C and 10.36% N), peach leaves (−25.95‰ $\delta^{13}C$ and 1.88‰ $\delta^{15}N$), and DORM (−17.27‰ $\delta^{13}C$ and 14.33‰ $\delta^{15}N$) were included in every run. Accuracy based on these standards is better than ±1.5% for C and N, ±0.4‰ for $\delta^{13}C$, and ±0.2‰ for $\delta^{15}N$.

Pollen and non-pollen palynomorphs (NPP) were identified in 1–2 cm^3 samples (*n*=102) that were treated with warm 10% KOH for 8 min, sieved through 150 μm mesh, and then treated with warm acetolysis for 2.5 min and mounted in 2000 cs silicone oil. Samples below 744 cm in the clay portion of the core were also treated with HF and sieved with 10 μm mesh. These five samples were excluded from NPP analysis because HF destroys many of these remains. One *Lycopodium* tablet of 18,584 ± 829 spores (Batch #177745) was added to each sample to estimate palynomorph concentrations. At least 400 terrestrial pollen and spores, not including *Sphagnum*, were identified in each sample. *Alnus* pollen were identified according to May and Lacourse (2012). Non-pollen palynomorphs including fungal spores, algal remains, aquatic plant microfossils and testate amoebae were identified using van Geel (1978), Pals et al. (1980), Charman et al. (2000), Clarke (2003) and Payne et al. (2012). Pollen percentages are based on all pollen and spores, except those from *Sphagnum* and obligate aquatic plants. Cluster analysis of the pollen percentage data was based on all taxa exceeding 5% of the sum except *Sphagnum* and aquatic taxa. Percentages were square root transformed and then analyzed using optimal splitting by information content and a broken stick model (Bennett, 1996). Cluster analysis of the NPP data was based on palynomorphs present in five or more samples using the same approach.

The >150 μm fraction of pollen samples was used for estimating peat composition following a quadrat technique similar to Barber et al. (1994). Each sample was poured into gridded Petri dishes and all remains were identified in 15 randomly selected 1×1 cm^2 quadrats.

Major peat components (herbaceous stems/leaves, moss stems/leaves, ligneous roots, ericad leaves, and unidentifiable organic material) were enumerated and are expressed as percentages of the total count in those quadrats. Other macrofossils (e.g., fungal sclerotia, *Nuphar* sclereids, charcoal) encountered in the same 15 quadrats were also noted.

RESULTS

Chronology, stratigraphy and peat composition

The age-depth model for the Grant's Bog core estimated an age of 13,316 cal yr BP (12,390–13,658 cal yr BP) for the base of the organic lake sediments at 744 cm (Fig. 2). Sediment and peat accumulation rates are typically between 0.03 and 0.08 cm/cal yr, although rates increase between 570 and 480 cm during accumulation of peat consisting mostly of herbaceous remains (Fig. 3), reaching a maximum of 0.24 cm/cal yr at 525 cm (8900 cal yr BP).

The base of the core (810–744 cm) consists of clay (Fig. 3). Simple wet mounts of these clays revealed marine diatoms (e.g. *Thalassiosira*, *Campylodiscus*) and *Dictyocha speculum* silicoflagellates below 765 cm, but both marine and freshwater algae (e.g. *Thalassiosira*, *Campylodiscus*, *Trachyneis aspera*, *Gyrosigma, Pediastrum*) between 765 and 744 cm. There is an abrupt transition at 744 cm from clay to lake sediment (744–693 cm) and then a gradual transition to limnic (possibly telmatic) peat by 693 cm. The limnic peat (693-628 cm) is composed of 40–65% herbaceous remains and 25–45% unidentifiable organic matter (UOM), but *Sphagnum* leaves and woody roots are also present (Fig. 3). *Nuphar* sclereids, likely derived from *N. polysepala*, are more abundant in this limnic peat than in the underlying lake sediment and are more or less absent above 618 cm. Peat consisting of 50–75% herbaceous remains and ~20% *Sphagnum* leaves occurs between 628 and 490 cm. *Scirpus* and *Dulichium arundinaceum* seeds and woody remains are also present and fungal sclerotia begin to appear more frequently (Fig. 3). Well-preserved *Sphagnum*-dominated peat occurs between 490 and 390 cm. This is overlain by mixed peat (390–234 cm) consisting of herbaceous, woody and *Sphagnum* remains as well as a higher amount of UOM. Ericaceae leaves, mycorrhizal roots and fungal sclerotia increase in this portion of the core. Peat dominated by herbaceous remains with abundant fungal sclerotia occurs from 234 to 78 cm. Peat near the surface (78–0 cm) is mixed in composition but marked by a notable increase in *Sphagnum* leaves. Macroscopic charcoal occurs throughout the core but is most abundant between 240 and 84 cm (Fig. 3).

A 1 mm tephra horizon is present at 280.5 cm. The age-depth model predicts an age of 5800 cal yr BP (5410–5970 cal yr BP) for this depth. This is within uncertainty of the age of the Glacier Peak-Dusty Creek tephra dated to 5120 ± 90 ^{14}C yr BP (5750–5940 cal yr BP) by Beget (1981) via charcoal embedded within a pyroclastic flow deposit near the base of Glacier Peak. Foit et al. (2004) report an interpolated age range of 5710–5880 cal yr BP for this tephra in lake sediments from southeastern British Columbia. We attempted to verify the identity of the tephra using electron microprobe analysis; however, the majority of the glass shards were too small for analysis with a 5 μm beam and only two returned quantitative results (Supplementary Material). For these two shards, a similarity coefficient of 0.93 shows that the glass composition matches Glacier Peak-Dusty Creek tephra better than all other Holocene tephras documented in southern British Columbia (Supplementary Material). The Glacier Peak-Dusty Creek tephra has not been recognized previously on Vancouver Island; however, Hansen (1950) hypothesized that tephra at depths of 2.8 and 3.0 m in a peat sequence at Black River Bog, ~5 km northwest of our study site, was derived from Glacier Peak. Deposition of the Glacier Peak-Dusty Creek tephra at

Grant's Bog, 350 km northwest of its source, represents its most northwesterly occurrence to date.

Bulk chemical and isotopic records

Changes in organic matter content (LOI) generally follow the overall stratigraphy. LOI increases rapidly from 3% in the basal clays to 30–70% in the overlying lake sediment (Fig. 4). The limnic peat is characterized by increasing LOI from 70 to 90%. LOI in the upper 6 m of terrestrial peat is 95–99%, although there is a minor decrease to 91% at 194 cm, immediately above large pieces of charcoal (0.5–1 cm^3) that were observed during subsampling. Ash-free bulk density (AFBD) is low in the basal clays (0.03 g/cm^3) and then increases gradually from 0.05 to 0.13 g/cm^3 between 744 and 194 cm, where it decreases abruptly to 0.07 g/cm^3. AFBD remains more or less at this lower density until 97 cm and then increases towards the surface.

Carbon and nitrogen also follow stratigraphic changes. Lake sediment at the base is about 30% C (Fig. 4). Carbon increases to 40% in the limnic peat and then to about 45% in the overlying terrestrial peat. Nitrogen is 2–3% in the lake sediment and limnic peat, and decreases gradually to about 1% in the terrestrial peat; however, there is a notable increase to 2–3% N between 3500–2300 cal yr BP. The C:N is <20 in the lake sediment and limnic peat (i.e., before 10,000 cal yr BP) and then increases gradually to 50–80 between 8700–3800 cal yr BP during accumulation of *Sphagnum* and mixed peat. Again, there is a notable decrease in C:N to 20–30 between 3500–2300 cal yr BP, before increasing to ~40 in the uppermost peat. $\delta^{13}C$ values are less than −29‰ in the lake sediment and limnic peat, and increase to about −27‰ in the terrestrial peat. $\delta^{15}N$ values are between −2‰ and 0‰ for much of the record. *Sphagnum* peat that accumulated 8700–7750 cal yr BP is marked by a decrease in $\delta^{15}N$ to −3.4‰.

Carbon accumulation rates (CAR) are generally low (5–10 g C/m^2/cal yr) in the lake sediment and limnic peat (Fig. 5), but then begin to increase dramatically at ~9700 cal yr BP during accumulation of peat consisting primarily of herbaceous remains, reaching a maximum of 81 g C/m^2/cal yr at 8900 cal yr BP. Carbon accumulation varies between 10 and 30 g C/m^2/cal yr for much of the mid-Holocene and then increases in the uppermost peat to ~40 g C/m^2/cal yr. Nitrogen accumulation rates (NAR) are ~0.5 g N/m^2/cal yr throughout most of the record (Fig. 5), but increase to 3 g N/m^2/cal yr at 8900 cal yr BP. Mean CAR and NAR, weighted by deposition time, are 19.5 g C/m^2/cal yr and 0.56 g N/m^2/cal yr, respectively, in the peat portion of the record. Time-weighted mean CAR for the various peat types are as follows: 8.3 g C/m^2/cal yr in the limnic peat; 38.6 and 16.3 g C/m^2/cal yr in the early and late Holocene herbaceous peat, respectively; 33.3 g C/m^2/cal yr in the mid-Holocene *Sphagnum* peat; and, 18.2 and 31.5 g C/m^2/cal yr in the mid- and late Holocene mixed peat, respectively.

Non-pollen palynomorphs

Cluster analysis identified five statistically significant non-pollen palynomorph (NPP) assemblage zones (Fig. 6; Supplementary Material) that generally follow changes in stratigraphy and C and N measurements. NPP assemblages in the lake sediment (13,300–11,600 cal yr BP) and limnic peat (11,600–9900 cal yr BP) are dominated by freshwater diatoms and *Filinia* type rotifer eggs, reflecting a freshwater environment. *Closterium* algae and fungal spores, including ascospores of *Kretzschmaria deusta* (a parasitic fungus on wood and roots), are generally more abundant in the limnic peat than the underlying lake sediment. NPP zone 2 (9800–8700 cal yr BP) coincides with accumulation of herbaceous peat and is marked by increases in *Closterium* algae, protists and Type 124 fungal spores. Assemblages in zone 3 (8700–4100 cal yr BP), which corresponds with accumulation of *Sphagnum* and mixed peat, are characterized primarily by *Assulina muscorum* and *Hyalosphenia subflava* protists, and *Entophlyctis lobata*, Microthyriaceae and *Gaeumannomyces* fungal remains. Fewer NPP types are present between 4100 and 2750 cal yr BP (zone 4), when %N increases (Fig. 4); however, there are notable increases in *Closterium* and Zygnemataceae algae and Type 124 fungal spores at this time (Fig. 6; Supplementary Material). Testate amoebae, particularly *H. subflava*, increase in the uppermost NPP zone 5. *Gelasinospora* and *E. lobata* fungal remains are also common. A number of testate amoebae including *Assulina muscorum*, *Arcella discoides* type, *Hyalosphenia papilio* and *Trigonopyxis arcula* type increase in the upper 12 cm of the core (Supplementary Material).

Pollen and spore assemblages

Cluster analysis identified four statistically significant pollen assemblage zones (Fig. 7). Pollen spectra between 14,000 and 13,300 cal yr BP, in the basal clays (zone 1), are 70–80% *Pinus contorta* type, 10% *Alnus viridis* type and up to 10% Cyperaceae. *Picea*, *Abies*, *Salix*, *Shepherdia canadensis*, Chenopodiaceae and Polypodiaceae are present in trace amounts. *Pinus contorta* continues to dominate the pollen record in the lake sediments of zone 2 (13,300–11,200 cal yr BP). *Abies* and *Picea* also increase, and *Pseudotsuga menziesii* and *Tsuga heterophylla* appear for the first time, although each of these account for less than 6%. *Alnus rubra* type increases abruptly to account for ~15% and *A. viridis* type remains at ~10%. Pollen from herbaceous plants and *Pteridium aquilinum* spores account for up to 4% and 7%, respectively. Aquatic taxa (*Typha*, *Nuphar polysepala*, *Brasenia schreberi*) are present in low relative abundance.

Pollen zone 3 (11,200–7800 cal yr BP) is marked by a dramatic decline in *Pinus contorta* to 25% and a corresponding increase in *Pseudotsuga menziesii* to 30–40% (Fig. 7). *Tsuga heterophylla* increases to 5–10% and *A. rubra* type accounts for 20–30%. There is an overall increase in non-arboreal pollen with Cyperaceae accounting for up to 7% and other herbaceous taxa including *Angelica* type and *Menyanthes trifoliata* accounting for another 3%. *Pteridium aquilinum* reaches its maximum abundance (12%) in zone 3. Pollen from aquatics is relatively abundant between 11,500 and 9700 cal yr BP during accumulation of limnic peat. Ericaceae pollen begins to increase at 9700 cal yr BP when terrestrial peat dominated by herbaceous remains begins accumulating. *Sphagnum* spores increase starting ~9500 cal yr BP, reaching 35–60% between 8900 and 7750 cal yr BP during accumulation of *Sphagnum*-dominated peat.

Pollen zone 4 (7800 cal yr BP to the present) is dominated by three main taxa: *Tsuga heterophylla* and *Alnus rubra* type, which are more abundant in subzone 4b, and *Pseudotsuga menziesii*, which is more abundant in subzone 4a (Fig. 7). *Pinus contorta* type accounts for 10–20% and Ericaceae increases relative to zone 3 with a few increases of up to 40%. There is also an isolated increase in *Sanguisorba* to 15% at 2800 cal yr BP. *Myrica* increases over the last 2700 cal yr but does not exceed 5%. *Pteridium aquilinum* accounts for 5–10% in subzone 4a and is present only intermittently in subzone 4b. *Sphagnum* spores are also generally more abundant in 4a than 4b. The uppermost samples are marked by a large increase in the relative abundance of *Alnus rubra*.

DISCUSSION

Wetland succession and C accumulation at Grant's Bog

The Grant's Bog core begins with marine clay deposited before 13,300 cal yr BP. This agrees with Hutchinson et al.'s (2004) sea level reconstruction based on isolation basins and ^{14}C-dated marine shells and wood in glaciomarine deposits that infers subaerial exposure of this area at 13,500 cal yr BP. Clay deposition occurred initially in a nearshore, marine environment and then in a brackish environment as relative sea level decreased.

A freshwater lake with *Nuphar polysepala* in low abundance and *Typha* and Cyperaceae at the margins was in place by 13,300 cal yr BP (Figs. 3 and 7), after the basin became isolated from marine waters. *Brasenia schreberi* was present in the lake by 12,700 cal yr BP (Fig. 7). Organic lake sediment with 2–3% N and a C:N less than 20 (Fig. 4), which is similar to most lakes (Meyers and Teranes, 2001), accumulated until 11,600 cal yr BP.

The gradual transition from organic lake sediment to limnic peat suggests the beginning of terrestrialization with decreasing lake levels and/or a reduction in lake area as well as increasing organic matter accumulation and potentially floating mat encroachment. Rotifer eggs, freshwater diatoms, *Typha* pollen, and *N. polysepala* pollen and sclereids are abundant in the limnic peat (Figs. 3, 6 and 7), suggesting the presence of standing water with emergent and floating-leaf aquatics until 9900 cal yr BP. Sclereids provide structural support and in Nymphaeaceae are more abundant in the petioles of erect aerial leaves than in floating lily pads (Etnier and Villani, 2007). The increase in *Nuphar* sclereids during accumulation of limnic peat (Fig. 3) likely reflects the presence of erect aerial leaves and/or an increase in the overall abundance of *N. polysepala* linked to decreasing water levels, since these aquatic plants tend to be most abundant in shallow-water wetlands and even occur in bog hollows on Vancouver Is.

today. *Angelica* type pollen (Fig. 7) suggests the nearby presence of *Angelica genuflexa*, a species characteristic of coastal BC wetlands. The limnic peat is also characterized by relatively high N (2–3%) and low $\delta^{13}C$ (< −30‰) due to the presence of aquatic plants and algal communities, which tend to be N-rich and ^{13}C-poor (Meyers and Teranes, 2001; Talbot, 2001). C and N accumulation are low during this marshy wetland stage (Fig. 5).

A small lake remains in the southeastern corner of the wetland complex today, but terrestrialization was more or less complete at the coring location by 9900 cal yr BP. $\delta^{13}C$ values become more positive in the warm, early Holocene (Fig. 4) and remain relatively high for the rest of the record, reflecting peat accumulation in a terrestrial setting (Jones et al., 2010; Andersson et al., 2012). Rapidly-accumulating herb-dominated peat with less UOM and a C:N increasing to 25–40 (Figs. 3 and 4) suggests a short-lived poor fen stage until ~8700 cal yr BP. Given the peat composition, the fen was likely dominated by sedges including *Dulichium arundinaceum* and *Scirpus* (likely *S. microcarpus*), but *Sphagnum* macro-remains and spores, ligneous roots, and Ericaceae, Cyperaceae and *Sanguisorba* pollen indicate diverse plant communities were present. The abundance of *Menyanthes trifoliata* pollen, *Archerella flavum* (syn. *Amphitrema flavum*) protists, and *Closterium* algae between 9700–8600 cal yr BP (Figs. 6 and 7) suggests wet conditions and a high water table.

Carbon and nitrogen accumulation increase dramatically during this fen stage, reaching maximum apparent rates of 81 g C/m^2/cal yr and 3 g N/m^2/cal yr, respectively, at 8900 cal yr BP (Fig. 5). These increases are linked to high peat accumulation rates (Fig. 2) combined with increasing bulk density (Fig. 4), as opposed to high C and N contents (Fig. 4). Since these high C and N accumulation rates are largely driven by a plateau in the age-model, the early Holocene increase is better described using time-weighted mean rates i.e., 48 g C/m^2/cal yr and 1.4 g

N/m^2/cal yr. An early Holocene increase in C and N accumulation was also found at other coastal BC peatlands (Turunen and Turunen, 2003; Lacourse and Davies, 2015) and is typical of northern peatlands (Loisel et al., 2014) including those in Alaska (e.g., Jones and Yu, 2010). These increases coincide with high summer insolation and the interval when seasonality in temperature is maximized (Fig. 5). Warm summers and increased seasonality favour peat accumulation by enhancing primary productivity during the growing season and reducing decomposition during winter (Asada and Warner, 2005; Yu et al., 2013; Loisel et al., 2014). Although the early Holocene was drier in coastal BC relative to modern (Walker and Pellatt, 2003; Brown et al., 2006), there was still sufficient moisture to promote peat accumulation and carbon storage (Gallego-Sala et al., 2018).

Carbon and nitrogen accumulation decrease with the development of a *Sphagnum*-dominated peatland ~8700 cal yr BP. At most peatlands in eastern North America, the transition from fen to bog occurred later, in the mid- to late Holocene (Yu et al., 2013). The well-preserved peat that accumulated at Grant's Bog ~8700–7750 cal yr BP is marked by an abundance of *Sphagnum* leaves (Fig. 3) and spores (Fig. 7). In general, fluctuations in the abundance of *Sphagnum* spores correlate well with changes in *Sphagnum* macro-remains, demonstrating that spores can provide a reliable record of *Sphagnum* abundance in some cases (cf. Lacourse and Davies, 2015). High C:N ratios of 60–80 and a notable decrease in $\delta^{15}N$ to −3.4‰ during this early Holocene *Sphagnum* phase (Fig. 4) suggest a low water table (Asada et al., 2005), reflecting the transition to ombrotrophy and a fully terrestrialized bog ecosystem. High concentrations of *Assulina muscorum*, a testate amoeba that is often most abundant in intermediate to dry peatlands (Charman et al., 2000; Payne et al., 2012), and remains from saprotrophic fungi (e.g., *Entophlyctis lobata* sporangia), which require oxic conditions to be major decomposers in peatlands, also suggest a lowering of relative water table depth compared to the preceding fen stage (Fig. 6). Ligneous roots that record colonization of the bog surface by woody plants are also present in this *Sphagnum* peat. Increases in Ericaceae leaf fragments (Fig. 3) and pollen (Fig. 7) suggest ericads were abundant in the plant community after 7750 cal yr BP and further isolation of the bog surface from the water table, despite increasing precipitation through the mid-Holocene (Brown et al., 2006).

Peat that accumulated in the mid-Holocene (7750–4700 cal yr BP) consists of a mixture of herbaceous, woody and *Sphagnum* remains with generally more UOM than before or after (Fig. 3). Increased decomposition in this portion of the record is also suggested by increases in mycorrhizal roots and fungal remains such as sclerotia, *E. lobata* sporangia, and *Gaeumannomyces* hyphopodia (Figs. 3 and 6). Isolated occurrences of *Drosera rotundifolia* pollen suggest nutrient-poor, acidic conditions. These various lines of evidence along with relatively high C:N of 40–60 (Fig. 4) indicate the site was an ombrotrophic peatland with mixed plant communities for much of the mid-Holocene. Carbon accumulation rates are only 15–20 g C/m^2/cal yr between 7200 and 1300 cal yr BP (Fig. 5); peat bulk density and C content (Fig. 4) are similar to the early Holocene, but accumulation rates (Fig. 2) are generally lower. Relative to the early Holocene, climate was cooler, wetter and less seasonal in the mid- and late Holocene (Walker and Pellatt, 2003; Brown et al., 2006; Lemmen and Lacourse, 2018). The abundance of macroscopic charcoal between ~4800 and 1000 cal yr BP (Fig. 3) suggests fire occurred on or near the peatland, despite a generally cooler, wetter climate.

Multiple proxies suggest changes in edaphic and hydrological conditions between 3500 and 2300 cal yr BP, likely as a result of disturbance by fire and subsequent flooding. The most striking change during this interval is an increase in N to 2–3% that is accompanied by a decrease in C:N to ~20 (Fig. 4), suggesting the surface of Grant's Bog was inundated. This interpretation is supported by coincident changes in NPP assemblages including increases in *Closterium* algae and diatoms (zone 4 in Fig. 6). The increase in Type 124 fungal spores, probably derived from *Persiciospora*, suggests eutrophic to mesotrophic conditions (Bakker and van Smeerdijk, 1982). A notable peak in *Sanguisorba* pollen and minor increases in *Salix* and Cyperaceae (Fig. 7), as well as a decrease in woody roots and ericad leaf fragments (Fig. 3), suggest a transition to plant communities more typical of fens than bogs in coastal BC. Large pieces of charcoal occur at a depth of 194–198 cm (~3600 cal yr BP), just before the increase in %N begins. Just above this, at 192–195 cm, there is a short-lived increase in ash content to 6–9% (Fig. 4), likely associated with an accumulation of combustion residues. Since a visible ash layer was not present, it is unlikely that the fire spread downwards to any great depth, as is the case in smouldering peat fires that tend to leave several cm of ash (Zaccone et al., 2014). Together, these various lines of evidence suggest the return to wet conditions and hydroseral reversion to a poor fen was initiated by fire, which would have enhanced nutrient availability and potentially altered local hydrology, rather than changes in climate.

Over the last 2000 cal yr, peat dominated by herbaceous remains graded into mixed peat with a C:N of ~40 (Fig. 4), indicating a return to drier surface conditions. *Hyalosphenia subflava* testate amoebae and *Gelasinospora* fungal spores, both of which are typical of dry conditions (Charman et al., 2000; Chambers et al., 2011; Payne et al., 2012), increase in the late Holocene (Fig. 6). *Myrica*, a nitrogen-fixing shrub typical of coastal wetlands, also increases (Fig. 7). Similar increases in *Myrica* occur in other pollen records from coastal BC (e.g., Mathewes, 1973; Brown and Hebda, 2002; Lacourse, 2005), suggesting a region-wide expansion of these shrubs in the late Holocene. At 1000 cal yr BP, there is a marked increase in *Sphagnum* leaves (Fig. 3) and spores (Fig. 7) with the uppermost 72 cm of peat composed of 30–80% *Sphagnum* and up to 35% ligneous roots. This is consistent with plant communities at Grant's Bog today, which are dominated by ericaceous shrubs (*Rhododendron groenlandicum*, *Kalmia microphylla* var. *occidentalis*, *Vaccinium uliginosum*) that tower above an almost complete moss cover of mostly *Sphagnum fuscum*, *S. angustifolium* and *S. capillifolium*. Testate amoebae concentrations, most notably *H. subflava*, increase in this well-preserved peat. In the uppermost 12 cm, which corresponds with peat in the acrotelm, a number of other testate amoebae typically found in nutrient-poor peatlands but with variable water table depths (Mitchell, 2004; Taylor et al., 2019) increase (e.g., *Assulina muscorum*, *Arcella discoides* type, *Hyalosphenia papilio*) or appear for the first time (e.g., *Trigonopyxis arcula* type, *Hyalosphenia elegans*) in the record (Fig. 6; Supplementary Material). Carbon and nitrogen accumulation rates increase over the last 1500 cal yr (Fig. 5), a trend found in most long-term records from northern peatlands (Loisel et al., 2014) that is in part explained by a shorter interval for decomposition loss following accumulation. Recent modelling efforts suggest C sequestration at mid- to high latitudes is likely to continue to increase through the 21st century with further warming (Gallego-Sala et al., 2018).

Postglacial Forest Dynamics near Grant's Bog

The pollen record from Grant's Bog provides insight into local changes in wetland vegetation but is primarily dominated by trees (Fig. 7), as would be expected given the high pollen productivity and effective pollen dispersal of conifers, which have dominated vegetation communities in the region since the late Pleistocene. There are few pollen records from central Vancouver Is. (Hansen, 1950; Heusser, 1960; Fitton, 2003; Mazzucchi, 2010) to compare to the lowland record from Grant's Bog. Most of these are from high elevations, have low sample resolution, and/or are poorly dated. In general, the record from Grant's Bog provides a history of compositional changes in forests that corresponds with expectations based on these and the many records to the south (e.g., Mathewes, 1973; Pellatt et al., 2001; Brown and Hebda, 2002, 2003; Gavin et al., 2013; Leopold et al., 2016) and north (e.g., Hebda, 1983; Lacourse, 2005; Galloway et al., 2007, 2009; Lacourse and Davies, 2015).

During and following the late-glacial decrease in relative sea level about 14,000 cal yr BP, vegetation near Grant's Bog consisted primarily of *Pinus contorta*, as was the case along much of the northeast Pacific coast at this time (Peteet, 1991; Lacourse, 2005; Lacourse et al., 2005; Galloway et al., 2009; Gavin et al., 2013; Leopold et al. 2016). *Alnus viridis*, *Salix* and *Shepherdia canadensis* shrubs were also present, making these early vegetation communities near Grant's Bog similar to those that occurred at about the same time at low elevations on northern Vancouver Is. (Lacourse, 2005) and the southeastern BC mainland (Mathewes, 1973). Macroscale climate at this time was cool and likely drier, relative to modern (Heusser et al., 1985; Kienast and MacKay, 2001; Lemmen and Lacourse, 2018).

Pinus contorta continued to dominate plant communities along much of the northeast Pacific coast until the beginning of the Holocene. At Grant's Bog, high relative abundance of *Pinus* pollen and moderate organic matter content suggest the presence of open *P. contorta* forests until about 11,200 cal yr BP (Figs. 4 and 7). Pollen from *Alnus rubra* and more-shade tolerant conifers including *Abies* and *Picea* increase during this interval, suggesting an increase in forest density at least regionally, as some portion of these probably derive from long-distance transport. *Pseudotsuga menziesii* first appears in the Grant's Bog record just before 13,000 cal yr BP. Given the short dispersal distance of *P. menziesii* pollen (Tsukada, 1982), it is likely that scattered individuals occurred near Grant's Bog by this time. *Pteridium aquilinum* ferns, which often occur in association with *P. menziesii* in modern forests, also appeared at this time.

The increase in *P. contorta* pollen in the Grant's Bog record between 12,700 and 11,700 cal yr BP is accompanied by minor increases in *Abies* and *Picea* and a decrease in *A. rubra* (Fig. 7). There is also a notable decrease in organic matter content (Fig. 4), indicative of lower overall productivity. The timing of these change suggests a link to Younger Dryas-related climate change. Some paleoecological records from the northeast Pacific coast suggest cooling during the Younger Dryas chronozone (Engstrom et al., 1990; Mathewes, 1993; Lacourse, 2005; Galloway et al., 2007; 2009; Gavin et al., 2013), while others show little evidence of cooling at this time (Brown and Hebda, 2003; Lacourse et al., 2005, 2012; Leopold et al., 2016). A chironomid-based temperature reconstruction from the south coast of BC suggests a decrease of as much as 3ºC, relative to modern, during the Younger Dryas (Lemmen and Lacourse, 2018), which is consistent with other paleotemperature records from the northeast Pacific (Kienast and MacKay, 2001; Gavin et al., 2013). It is unlikely that the Younger Dryas on the northeast Pacific coast was cool and dry, as it was at many locations around the North Atlantic and in northeast Asia (Björck, 2007). Some of the strongest terrestrial evidence for cooling in the northeast Pacific is an increase in *Tsuga mertensiana* (Mathewes, 1993; Lacourse, 2005), an indicator of cool and moist climate. Recent modelling efforts suggest the Younger Dryas chronozone in the northeast Pacific was characterized by an increase in moisture (Renssen et al., 2018).

At ~11,200 cal yr BP, there was a rapid transition to *Pseudotsuga menziesii* forest with abundant *Pteridium aquilinum* ferns in the understorey. *Pseudotsuga* continued to dominate forests near Grant's Bog until about 7000 cal yr BP. The expansion of *P. menziesii* populations is a well-documented feature of Holocene vegetation change along the south coast of BC. During the last glacial maximum, *P. menziesii* occurred south of the Cordilleran Ice Sheet in western Washington and Oregon (Tsukada, 1982; Gugger and Sugita, 2010). Populations migrated north as the ice sheet receded, reaching southern Vancouver Is. as early as 14,000 cal yr BP (Brown and Hebda, 2003), east-central Vancouver Is. by 13,000 cal yr BP (this study), and northern Vancouver Is. by ~11,000 cal yr BP (Lacourse, 2005; Lacourse and Davies, 2015), at an approximate rate of 120 m/yr. Early Holocene warming allowed *P. menziesii* to become the dominant conifer on southern (Pellatt et al., 2001) and central (this study) Vancouver Is. by 11,000 cal yr BP. It continues to be abundant near Grant's Bog today and to the south, but it has been uncommon on northern Vancouver Is. since about 7500 cal yr BP (Lacourse, 2005; Lacourse and Davies, 2015), when cooler and moister climate lead to contraction of its northern limit (Gugger and Sugita, 2010).

Alnus rubra and *Tsuga heterophylla* pollen increase at Grant's Bog in the early Holocene, but this is probably a reflection of their increasing populations throughout the region, rather than an indication of high local abundance. Only a few Cupressaceae pollen grains, likely from *Thuja plicata*, are present between 11,500 and 8000 cal yr BP (Fig. 7). Cupressaceae is otherwise absent from the Grant's Bog record. This is unusual compared to Holocene pollen records to the south (Pellatt et al., 2001; Lucas and Lacourse, 2013; Leopold et al., 2016) and north (Lacourse, 2005; Galloway et al., 2007), which contain small, but consistent amounts of Cupressaceae in the early Holocene. Furthermore, most pollen records show increasing relative abundance of Cupressaceae in the mid- to late Holocene as precipitation increased across the region (Brown et al., 2006). However, there are a few records, mostly from eastern Vancouver Is. in the rainshadow of the Vancouver Island Ranges (Brown and Hebda, 2003; Mazzucchi, 2010; Lacourse and Davies, 2015), where Cupressaceae pollen is infrequent throughout the Holocene.

Thuja plicata is generally intolerant of dry conditions and is most abundant in wet coniferous forests, although it can also occur in drier *P. menziesii* forests, at least on moister sites. Allen et al. (1999) found little Cupressaceae pollen in modern surface samples from *P. menziesii* forests, except at moist sites close to *Tsuga heterophylla*-dominated forest. The low relative abundance of Cupressaceae in the Grant's Bog record suggests it was not abundant near the site. Rainshadow effects including dry summers in particular likely created conditions with insufficient moisture to support *Thuja plicata*.

Increasing precipitation and decreasing temperature in the mid- and late Holocene (Walker and Pellatt, 2003; Brown et al., 2006) facilitated the expansion of *Tsuga heterophylla* near Grant's Bog and throughout much of the region. Mid-Holocene forests near Grant's Bog were composed primarily of *P. menziesii* and *T. heterophylla* with *P. aquilinum* ferns continuing in the understorey. By 5000 cal yr BP, the abundance of *T. heterophylla* and *Alnus rubra* increased further and *P. menziesii* and *P. aquilinum* decreased, which suggests an increase in forest density and relatively closed canopies in these *T. heterophylla*–*P. menziesii* forests. The uppermost pollen assemblages at Grant's Bog are characterized by an increase in *A. rubra* to 60−75%. A similar increase was found at Port McNeill Bog (Lacourse and Davies, 2015), approximately 160 km to the northwest of Grant's Bog. Given this species tendency to colonize disturbed sites, it is likely that at least some portion of this increase is linked to increased human activity and logging in the region.

CONCLUSIONS

A complete terrestrialization sequence is recorded at Grant's Bog, starting with isolation of the basin from marine waters by 13,300 cal yr BP due to decreases in relative sea level. Transition from an oligotrophic lake to a shallow, marshy wetland with aquatic plants occurred by ~11,600 cal yr BP. This was followed by autogenic development of a poor fen by 9900 cal yr BP and then a drier ombrotrophic bog by 8700 cal yr BP. Changes in multiple paleoenvironmental proxies since the mid-Holocene point towards fluctuating edaphic and hydrological conditions and local plant community composition, including hydroseral reversion to a poor fen 3500–2300 cal yr BP following disturbance by fire.

Long-term mean C accumulation rates at peatlands on the northeast Pacific coast (this study; Turunen and Turunen, 2003; Jones and Yu, 2010; Nichols et al., 2014; Lacourse and Davies, 2015) are low compared to some continental peatlands (Yu et al., 2014; Zhao et al., 2014), suggesting that seasonality plays an important role in constraining peatland C sequestration. Mild year-round temperatures on the coast lead to long growing seasons and enhanced primary productivity, but also to greater decomposition and lower net C storage compared to inland peatlands with more seasonal climates (Asada and Warner, 2005). This pattern is also true on long timescales at many sites including Grant's Bog, where the highest C accumulation rates occurred during the early Holocene when summers were warmer and seasonality was maximized. Carbon accumulation was high in the early Holocene but comparatively low in the cooler, less seasonal late Holocene, despite accumulation of peat that is similar in composition. Macroscale climate appears to be a more dominant control on long-term C accumulation in peatlands than local, site-specific factors such as peat composition or vegetation type. Further research comparing a large number of sites is needed to confirm this general pattern.

The pollen record from Grant's Bog indicates that open *Pinus contorta*-dominated communities were present by 14,000 cal yr. *Pseudotsuga menziesii* forest with abundant

Pteridium aquilinum ferns were established in the early Holocene and dominated coastal lowlands until ~7000 cal yr BP. *Tsuga heterophylla* and *P. menziesii* formed closed-canopy forests in the mid-Holocene. In contrast to most other pollen records from coastal BC, Cupressaceae (*Thuja plicata*) appears never to have been abundant in forests near Grant's Bog. Additional pollen records from eastern Vancouver Island should clarify the spatial extent of this pattern.

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Tables

Table 1. AMS radiocarbon and calibrated ages from Grant's Bog on Vancouver Island, British Columbia.

Figure Captions

Figure 1. Map of Vancouver Island on the south coast of British Columbia, Canada showing the location of Grant's Bog (star) and other paleoecological studies mentioned in the text: 1 – Two Frog Lake and Tiny L. (Galloway et al., 2007, 2009), 2 – Bear Cove Bog (Hebda, 1983), 3 – Misty L. (Lacourse, 2005), 4 – Port McNeill Bog (Lacourse and Davies, 2015), 5 – Harris Lake Ridge Bog (Fitton, 2003) and Burman Pond (Mazzucchi, 2010), 6 – Turtle L. (Fitton, 2003); 7 – Marion L. (Mathewes, 1973); 8 – Porphyry L. (Brown and Hebda, 2003); 9 – Roe L. (Lucas and Lacourse, 2013), 10 – Saanich Inlet (Pellatt et al., 2001), 11 – Killebrew L. (Leopold et al., 2016), 12 – East Sooke Fen (Brown and Hebda, 2002). Top inset shows location in North America. Bottom inset shows monthly means of minimum and maximum temperature and precipitation at nearby Black Creek climate station (Environment Canada, 2018).

Figure 2. Age-depth model for the Grant's Bog core from Vancouver Is., British Columbia. Grey bands are 95% confidence intervals based on 10,000 model runs. The age at 727 cm was excluded from the model. Glacier Peak-Dusty Creek tephra was not used to constrain the age model; its depth and modelled age of 5800 cal yr BP are shown with dashed droplines.

Figure 3. Stratigraphy, peat components and plant macrofossils for the core from Grant's Bog on Vancouver Is., British Columbia. Circles indicate depth of infrequent macrofossils. Note that the

charcoal scale is truncated. Glacier Peak-Dusty Creek tephra is shown as a horizontal line at 280 cm in the stratigraphic column. Triangles along the y-axis show the position of ^{14}C ages (Table 1) used to build the age model. UOM=Unidentifiable organic matter.

Figure 4. Stratigraphy and bulk chemical and isotopic records for Grant's Bog on Vancouver Is., British Columbia. See. Fig. 3 for stratigraphy legend. AFBD=Ash-free bulk density.

Figure 5. Carbon and nitrogen accumulation rates at Grant's Bog on Vancouver Is., British Columbia. Mean accumulation rates in 500 cal yr bins are overlaid on apparent rates. Also shown are % wet (black) and dry (grey) peat components from Fig. 3, C:N, and January and July insolation anomaly at 50°N (Berger and Loutre, 1991). Grey box highlights the period of maximum difference between January and July insolation.

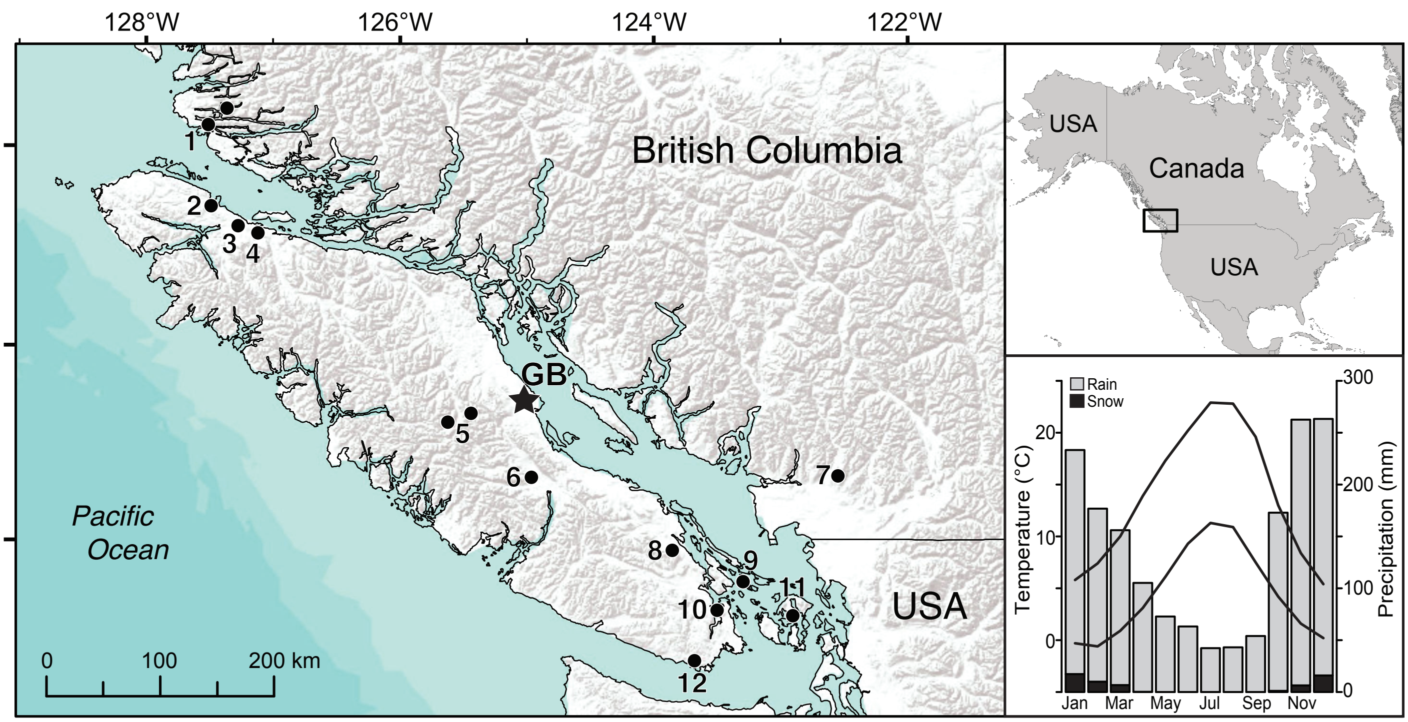
Figure 6. Concentrations of abundant non-pollen palynomorphs at Grant's Bog on Vancouver Is., British Columbia. Other testate amoebae are mostly *Arcella discoides* type and *Hyalosphenia papilio* but also include *Trigonopyxis arcula* type, *Assulina seminulum* and *H. elegans*. Note changes in scale on x-axes. Numbers in parentheses are NPP types in van Geel (1978) and Pals et al. (1980). See Fig. 3 for stratigraphy legend.

Figure 7. Pollen and spore percentages from Grant's Bog on Vancouver Is., British Columbia with 5× exaggeration (grey silhouettes) for infrequent taxa. Ericaceae total includes *Empetrum nigrum*, *Ledum* and *Vaccinium*, but is mostly undifferentiated Ericaceae pollen. Apiaceae is

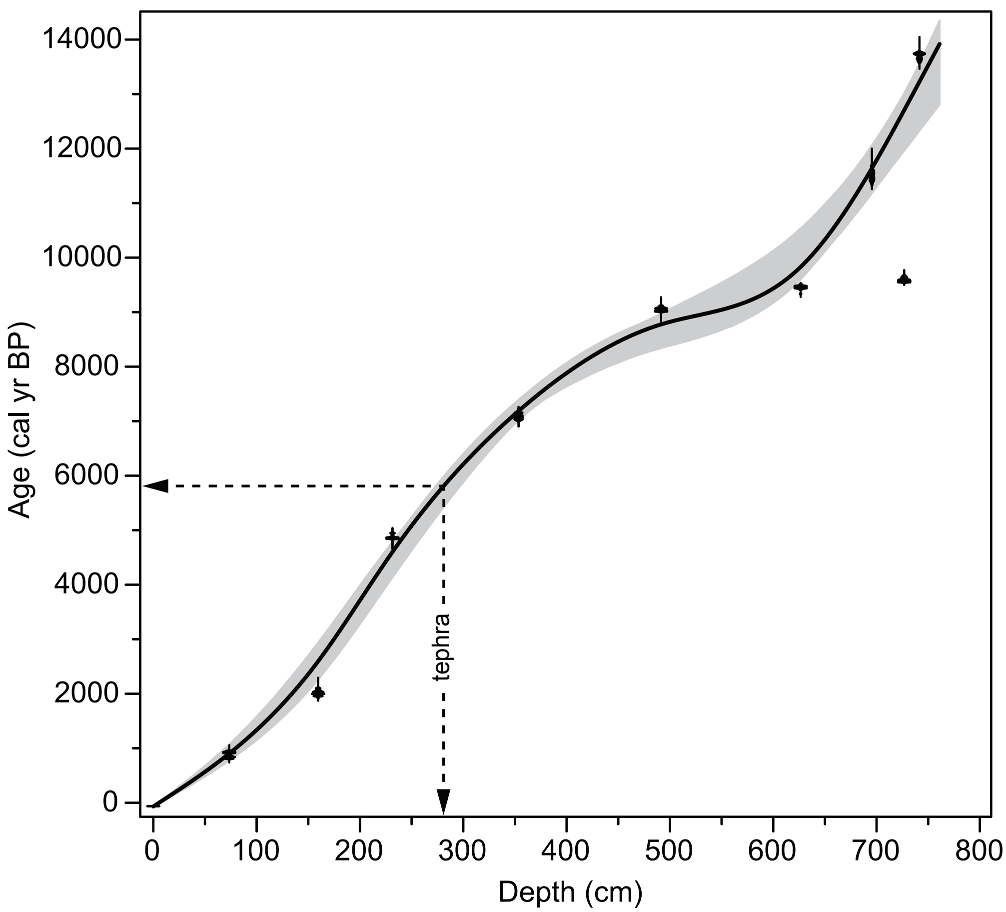
almost exclusively *Angelica* type. Order of taxa is based on weighted averages. Grey horizontal band marks the Younger Dryas chronozone.

Table 1. AMS radiocarbon and calibrated ages from Grant's Bog on Vancouver Island, British Columbia.

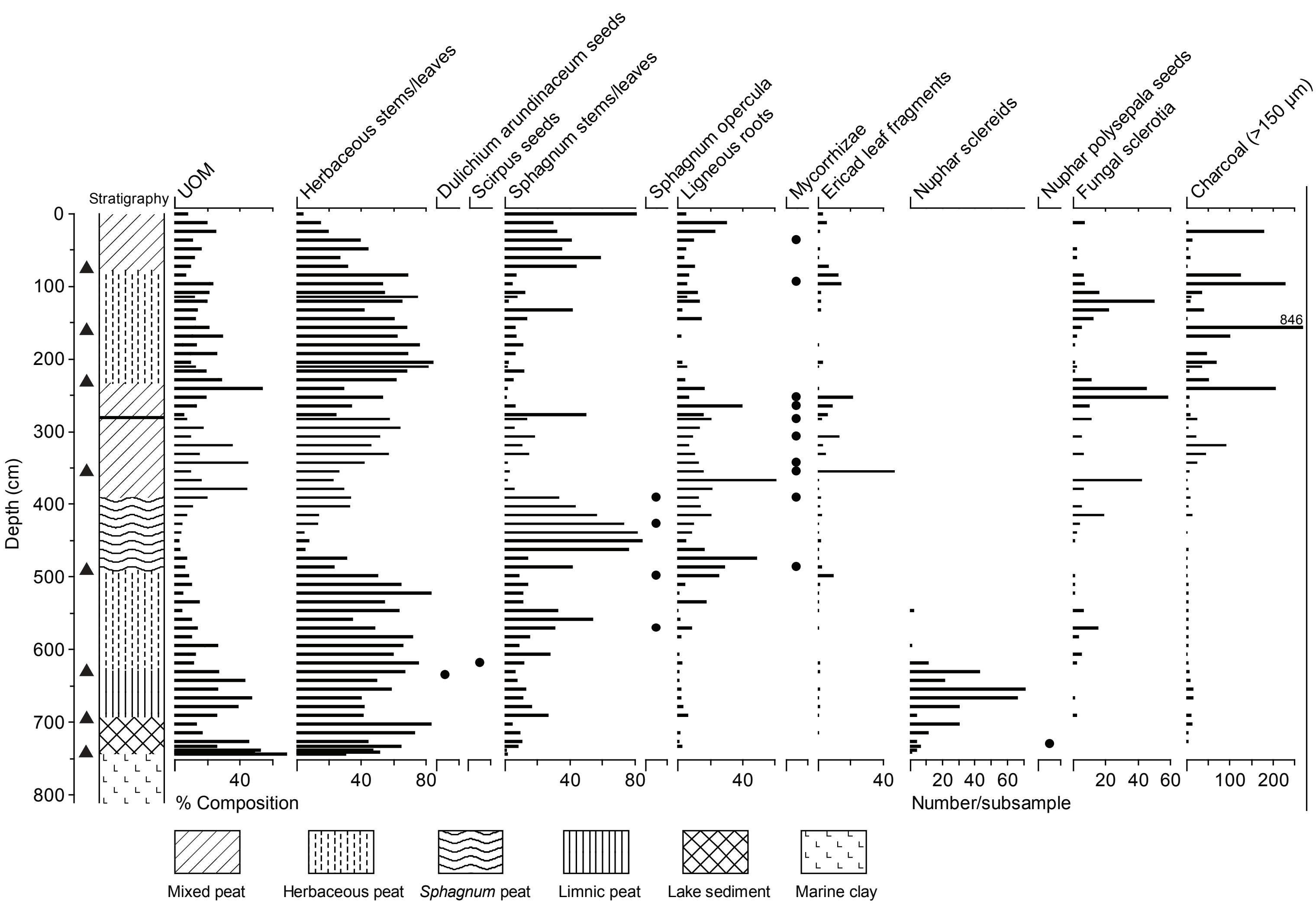
Depth (cm)	Material	$\delta^{13}C$ (‰)	Radiocarbon Age (^{14}C yr BP ± 1σ)	2σ Calendar Age Range (cal yr BP)	Lab Code
73–74	Wood	–25.7	990 ± 30	800 – 960	Beta-439738
159–160	Plant macrofossils (>180 μm) extracted from herbaceous peat	–26.4	2050 ± 30	1930 – 2110	Beta-463068
231–232	Wood	–23.6	4300 ± 30	4830 – 4960	Beta-439739
353–354	Plant macrofossils (>180 μm) extracted from ericad-herbaceous peat	–28.7	6190 ± 30	6990 – 7170	Beta-463069
491–492	Wood	–28.2	8110 ± 30	9000 – 9120	Beta-439740
626–627	Wood	–27.2	8420 ± 30	9410 – 9520	Beta-439741
695–696	Organic lake sediment	–32.5	10,020 ± 30	11,330 – 11,700	Beta-480750
726.5–727	Wood	–27.8	8640 ± 30	9540 – 9670	Beta-439742
741–742	Organic lake sediment	–29.7	11,900 ± 40	13,570 – 13,790	Beta-475650



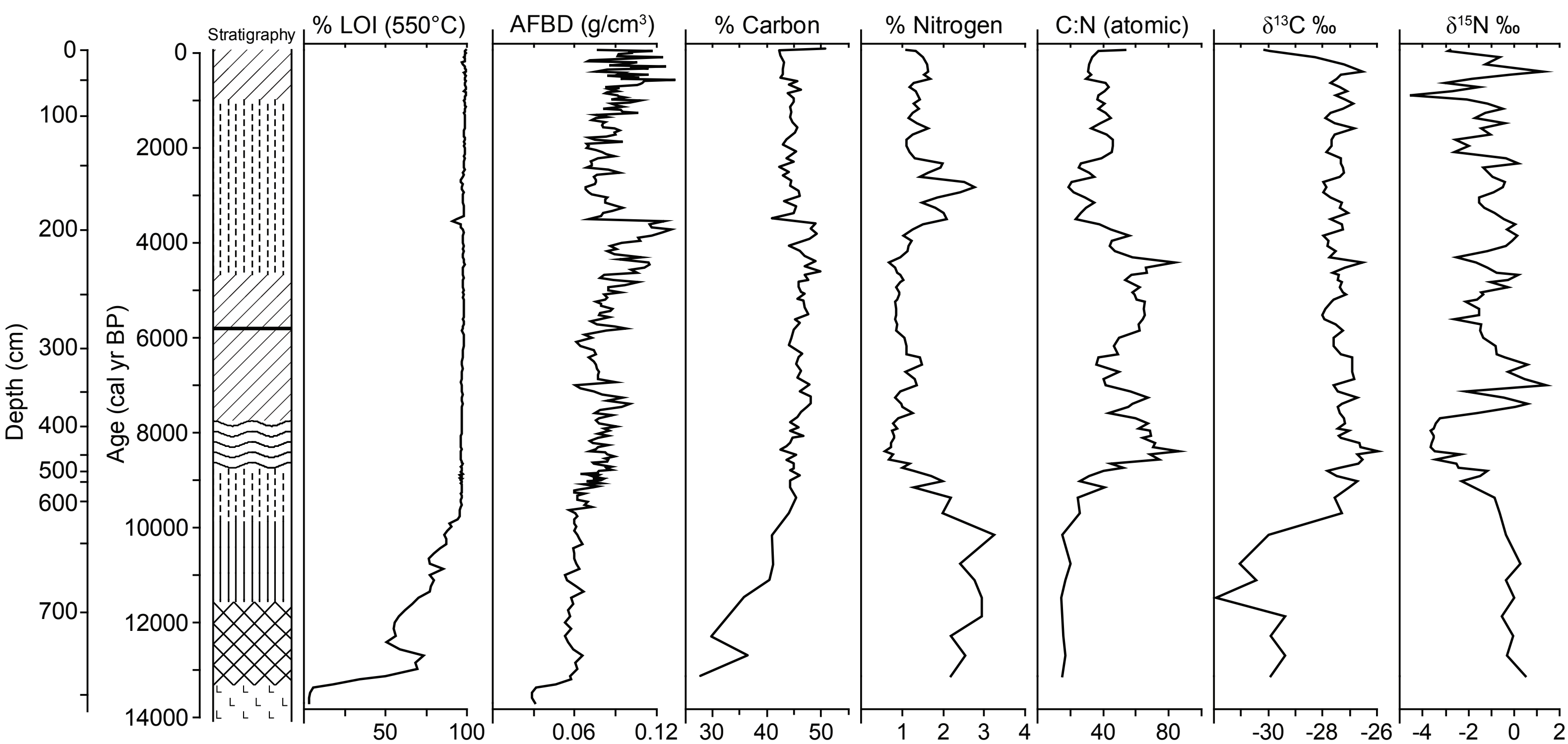
128°W
126°W
124°W
122°W
51°N
50°N
49°N
British Columbia
Pacific Ocean
USA
GB
1
2
3
4
5
6
7
8
9
10
11
12
0
100
200 km
USA
Canada
USA
Rain
Snow
Temperature (°C)
Precipitation (mm)
Jan
Mar
May
Jul
Sep
Nov



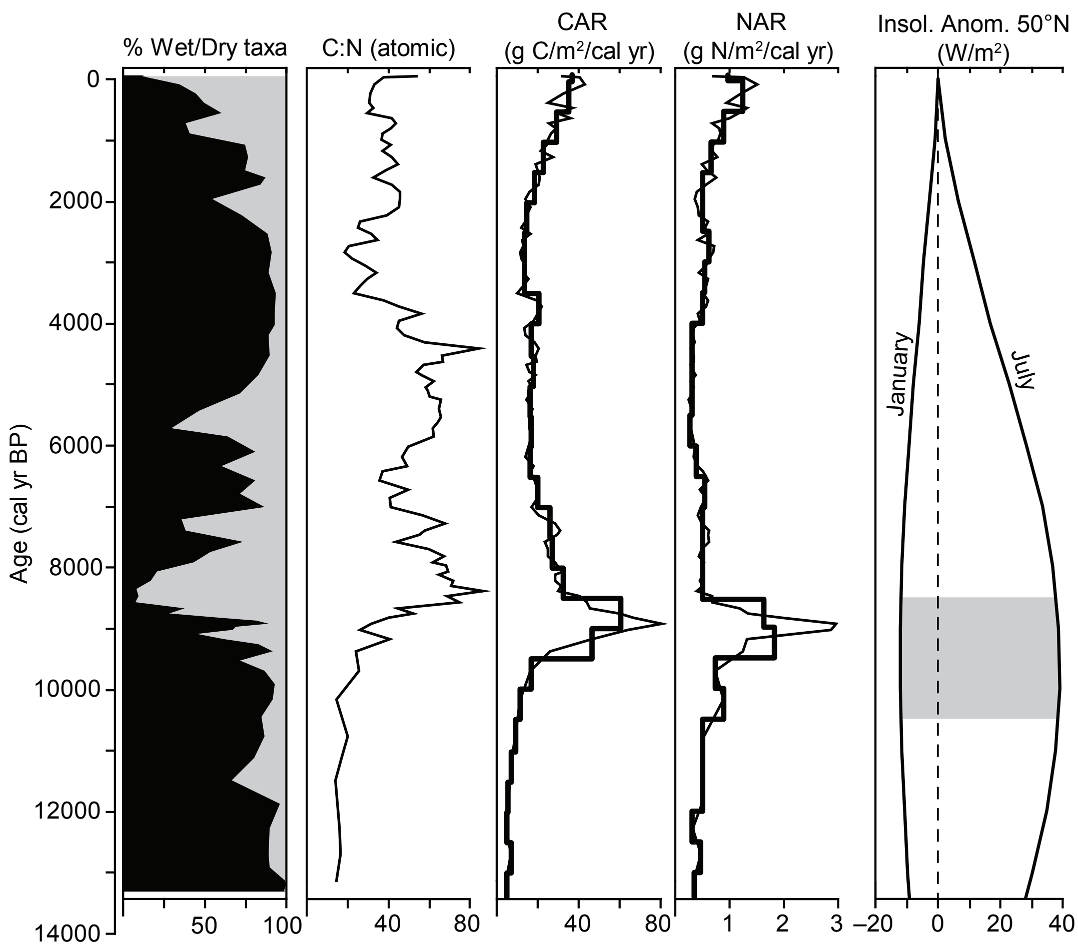
14000
12000
10000
8000
6000
4000
2000
0
Age (cal yr BP)
tephra
0
100
200
300
400
500
600
700
800
Depth (cm)



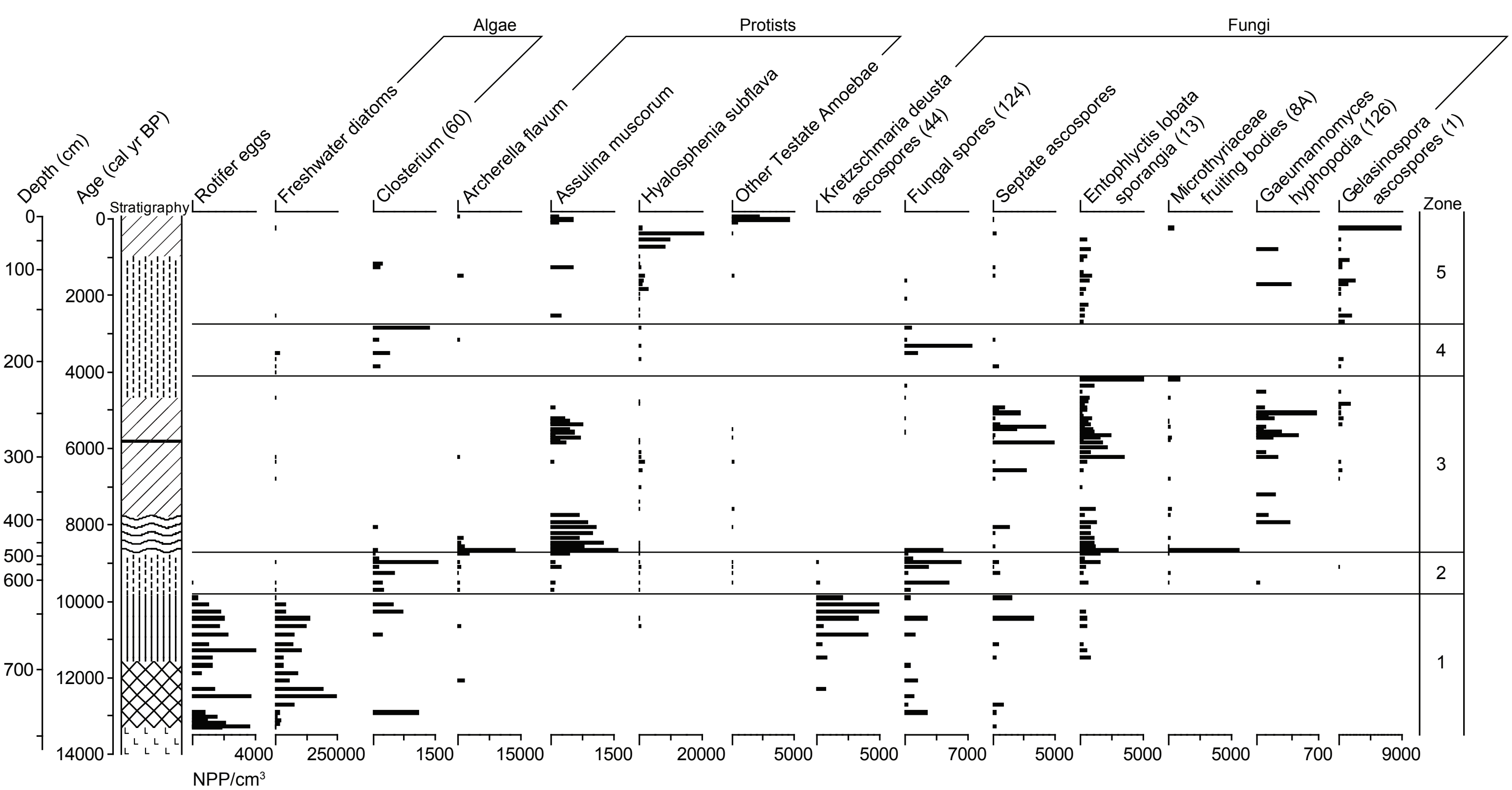
Stratigraphy
UOM
Herbaceous stems/leaves
Dulichium arundinaceum seeds
Scirpus seeds
Sphagnum stems/leaves
Sphagnum opercula
Ligneous roots
Mycorrhizae
Ericad leaf fragments
Nuphar sclereids
Nuphar polysepala seeds
Fungal sclerotia
Charcoal (>150 µm)
Depth (cm)
% Composition
Number/subsample
846
Mixed peat
Herbaceous peat
Sphagnum peat
Limnic peat
Lake sediment
Marine clay



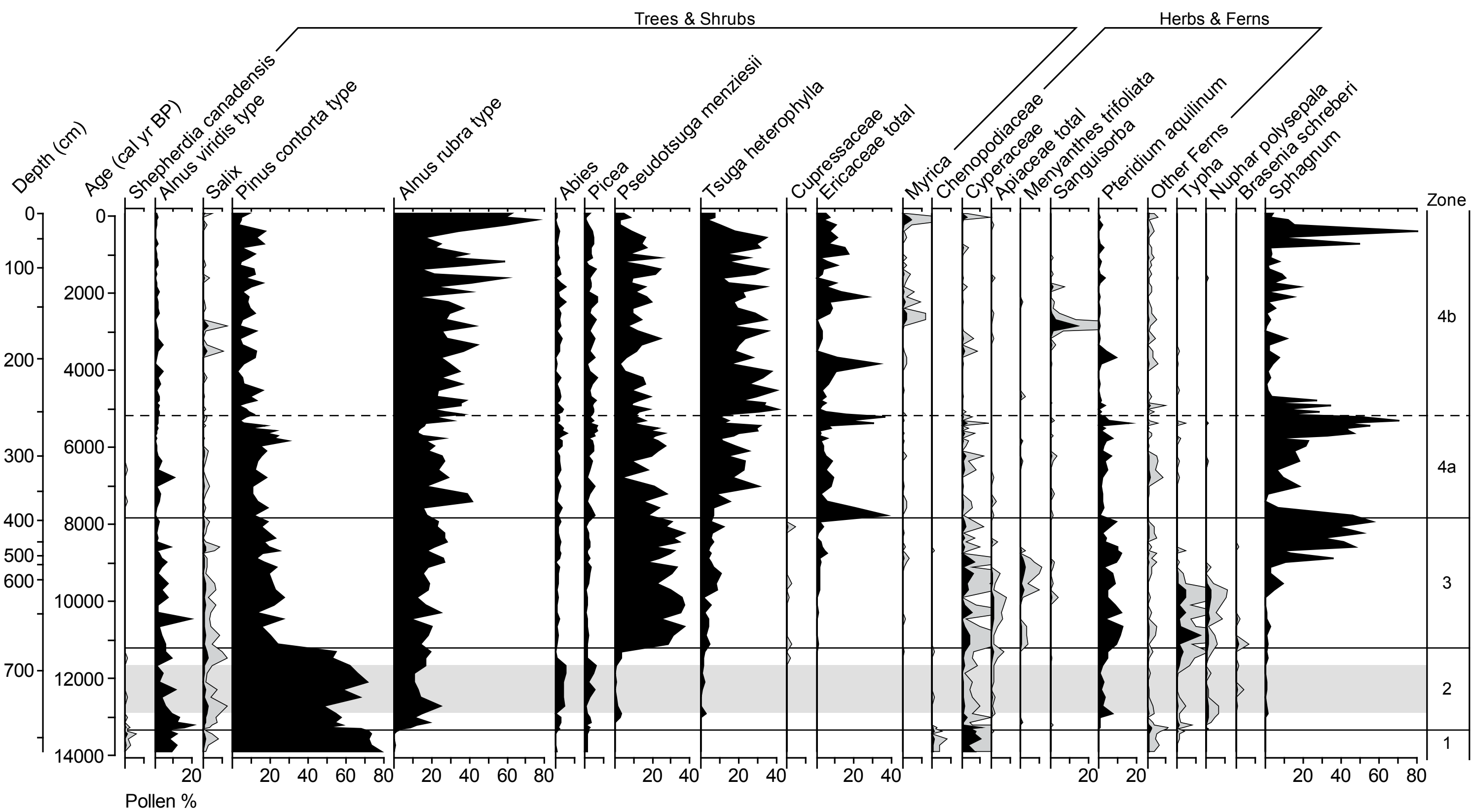
Depth (cm)
0
100
200
300
400
500
600
700
Age (cal yr BP)
0
2000
4000
6000
8000
10000
12000
14000
Stratigraphy
% LOI (550°C)
50
100
AFBD (g/cm³)
0.06
0.12
% Carbon
30
40
50
% Nitrogen
1
2
3
4
C:N (atomic)
40
80
δ¹³C ‰
-30
-28
-26
δ¹⁵N ‰
-4
-2
0
2



% Wet/Dry taxa
C:N (atomic)
CAR
(g C/m²/cal yr)
NAR
(g N/m²/cal yr)
Insol. Anom. 50°N
(W/m²)
Age (cal yr BP)
0
2000
4000
6000
8000
10000
12000
14000
50
100
40
80
40
80
1
2
3
–20
0
20
40
January
July

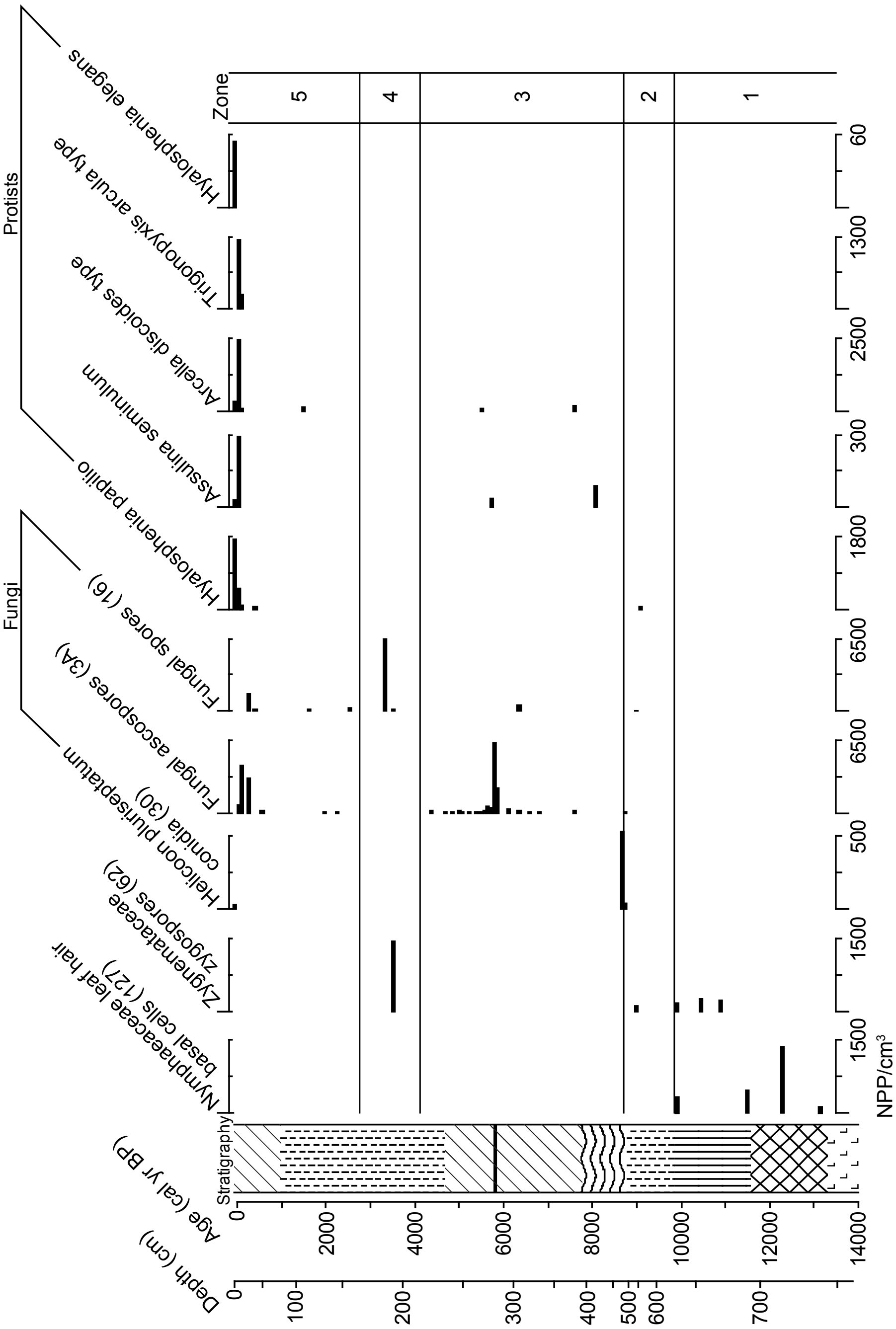


Algae
Protists
Fungi
Depth (cm)
Age (cal yr BP)
Stratigraphy
Rotifer eggs
Freshwater diatoms
Closterium (60)
Archerella flavum
Assulina muscorum
Hyalosphenia subflava
Other Testate Amoebae
Kretzschmaria deusta ascospores (44)
Fungal spores (124)
Septate ascospores
Entophlyctis lobata sporangia (13)
Microthyriaceae fruiting bodies (8A)
Gaeumannomyces hyphopodia (126)
Gelasinospora ascospores (1)
Zone
0
100
200
300
400
500
600
700
2000
4000
6000
8000
10000
12000
14000
4000
250000
1500
15000
1500
20000
5000
5000
7000
5000
5000
5000
700
9000
NPP/cm³
1
2
3
4
5



Trees & Shrubs
Herbs & Ferns
Depth (cm)
Age (cal yr BP)
Shepherdia canadensis
Alnus viridis type
Salix
Pinus contorta type
Alnus rubra type
Abies
Picea
Pseudotsuga menziesii
Tsuga heterophylla
Cupressaceae
Ericaceae total
Myrica
Chenopodiaceae
Cyperaceae
Apiaceae total
Menyanthes trifoliata
Sanguisorba
Pteridium aquilinum
Other Ferns
Typha
Nuphar polysepala
Brasenia schreberi
Sphagnum
Zone
4b
4a
3
2
1
Pollen %

Supplementary Material for Lacourse et al. 2019. ***Postglacial wetland succession, carbon accumulation and forest dynamics on the east coast of Vancouver Island, British Columbia, Canada*****. Quaternary Research.**



Supplementary Figure 1: Concentrations of infrequent non-pollen palynomorphs in the core from Grant's Bog on Vancouver Island, British Columbia. Note changes in scale on x-axes. Numbers in parentheses are NPP types in van Geel (1978) and Pals et al. (1980). See Fig. 3 of the paper for stratigraphy legend.

Electron microprobe analysis of tephra from Grant's Bog, Vancouver Island

Glass shards from peat at Grant's Bog were mounted in epoxy on a glass slide and polished to 0.25-μm diamond grit. We attempted to determine major element concentrations by electron microprobe analysis at the University of British Columbia. Operating conditions were 15 kV and 25 nA, with a 5 μm beam, with peak counting times as follows: Si, Al, K (15 s); Ti, Mn, Fe (25 s); Mg, Ca (20 sec); and Na (4 s). To limit volatilization by the electron beam, Na and K were counted first, and background counts were only made on first and last samples in the analytical run. Fragments of a Mt. Meager dacite glass were also analyzed to check for accuracy and showed results within 5% of values reported in Rust et al. (1999) for all elements except Na (20%) and Ti (25%), which occur at trace levels. The majority of the glass shards from Grant's Bog were too small for analysis with a 5 μm beam, and only 2 of 18 glass shards returned quantitative results. Nonetheless, we compared the mean composition of these two shards to nine other tephras known to occur in southern British Columbia as well as two mid-Holocene tephras from Alaska. Only the tephra from Mount Mazama has ever been documented on Vancouver Island.

Glass shards at Grant's Bog are relatively high in SiO_2 and K_2O, and low in Al_2O_3 and Na_2O, making them similar to tephras from Glacier Peak but dissimilar to those from other nearby possible sources (Supplementary Table 1 and Supplementary Figure 2). The highest similarity coefficient (0.93) is with the Glacier Peak-Dusty Creek tephra that has been dated to 5120 ± 90 ^{14}C yr BP (5750−5940 cal yr BP) by Beget (1981) via charcoal embedded in pyroclastic flow near the base of Glacier Peak. Foit et al. (2004) report an interpolated age range of 5710−5880 cal yr BP for this tephra in lake sediments from southeastern British Columbia. The age for the tephra based on the Grant's Bog age-depth model (5800 cal yr BP) is within the uncertainty of both of these age estimates and provides further support that the tephra at Grant's Bog is derived from the Glacier Peak-Dusty Creek assemblage.

Supplementary Table 1: Major element composition of glass shards (normalized to 100 wt. %) from Grant's Bog, Vancouver Island and various standards reported in the literature. The age-depth model for Grant's Bog predicts an age of 5800 cal yr BP (5410–5970 cal yr BP) for the depth of the tephra.

Shard	**SiO_2**	**TiO_2**	**Al_2O_3**	**Fe_2O_3**	**MgO**	**CaO**	**Na_2O**	**K_2O**	**Cl**	**MnO**	**Total**	**Analytical Total**
GB280-1	78.40	0.19	12.17	0.95	0.14	1.08	2.83	3.89	0.35	0.003	100	85.6
GB280-2	79.95	0.30	10.90	1.03	0.15	0.89	2.28	4.29	0.14	0.059	100	93.3
Mean	79.18	0.24	11.53	0.99	0.15	0.99	2.56	4.09	0.24	0.03	100	
Std. Dev.	1.10	0.08	0.90	0.05	0.01	0.13	0.39	0.29	0.15	0.04		

Tephra	**SiO_2**	**TiO_2**	**Al_2O_3**	**Fe_2O_3**	**MgO**	**CaO**	**Na_2O**	**K_2O**	**SC [a]**	**Age (cal yr BP)**
GP DC [b]	78.24	0.19	12.15	1.03	0.16	0.93	3.51	3.65	0.93	5750–5940
GP A [c]	77.35	0.38	12.32	1.41	0.18	0.86	3.28	4.10	0.88	~1700–2000
GP D [d]	78.16	0.38	11.68	1.35	0.13	0.71	3.42	4.10	0.87	~6000–6300
GP G [e]	77.41	0.18	12.73	1.03	0.23	1.21	3.62	3.39	0.87	13,410–13,710
GP B [e]	77.23	0.21	12.78	1.15	0.27	1.41	3.72	3.00	0.82	13,410–13,710
Oshetna [f]	72.55	0.43	14.63	2.29	0.66	2.31	4.20	2.72	0.80	~6500–7000
Aniakchak [g]	70.86	0.48	15.00	2.70	0.54	1.73	5.77	3.03	0.79	~3600
Meager [h]	75.21	0.34	13.52	1.59	0.35	1.28	4.51	3.25	0.77	2300–2400
MSH-P [i]	76.90	0.21	13.11	1.61	0.32	1.18	3.91	2.65	0.77	~2700–3000
MSH-Y [i]	76.40	0.14	13.71	1.37	0.31	1.56	4.21	2.22	0.72	~3400–3700
Mazama [j]	73.26	0.42	14.34	2.26	0.43	1.59	4.80	2.74	0.66	7580–7680

[a] Similarity coefficient (SC) is the weighted average of the ratios between the mean composition of the Grant's Bog tephra and standards, after Borchardt et al. (1972), using all oxide concentrations except Cl and MnO. As per Foit et al. (2014), TiO_2 and MgO are weighted to 0.25 because of low concentrations and high relative error. Na_2O is also weighted to 0.25 because of volatilization of sodium that typically occurs during analysis of glass. A SC of 1 represents a perfect match.

[b] Glacier Peak-Dusty Creek, north-western Washington: Hallett et al. (2001), Beget (1981, 1984)

[c] Glacier Peak A, north-western Washington: Foit et al. (2004), Mastin and Waitt (2000)

[d] Glacier Peak D, north-western Washington: Foit et al. (2004), Beget (1984)

[e] Glacier Peak G and B, north-western Washington: Kuehn et al. (2009)

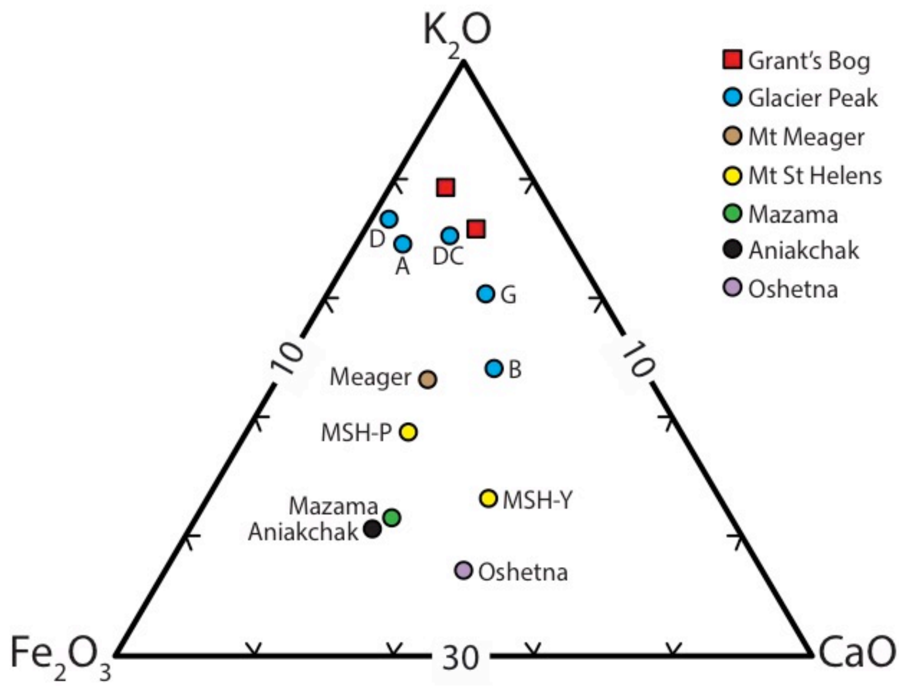
[f] Oshetna, south-central Alaska: Child et al. (1998), Dixon and Smith (1990)

[g] Aniakchak, south-western Alaska: Denton and Pearce (2008), Beget et al. (1992)

[h] Mount Meager Bridge River UA1 (Pebble Creek), south-western British Columbia: Westgate (1977), Leonard (1995)

[i] Mount St. Helens P and Y, south-western Washington: Foit et al. (2004), Mullineaux (1996)

[j] Mount Mazama, south-western Oregon: Foit et al. (1993), Egan et al. (2015)



Supplementary Figure 2: Relative abundance of K_2O-Fe_2O_3-CaO in tephras given in Supplementary Table 1. Red squares are shards from Grant's Bog. Circles are tephra standards. Tephras from Glacier Peak (blue circles) are labelled with letters as per Supplementary Table 1.

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